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- The seismic reflection takes an interest essentially to the P compression waves, in the setting of seismic studies for the geo-technical;
- the recording of the S waves can to prove out to be useful for the shearing modules calculation.
- In knowing the velocity of the P waves and S, it is possible of calculating the Poisson's ratio: 48.

BASIC THEORETICAL PRINCIPLES OF SEISMIC METHODS I. Huygens ...

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$K = \frac{1}{\rho v}$ : The 19. Figure 2.1: Particle motion of P and S waves solution to this equation is:  $p(x;t) = s(t - x/c)$  (2.4) where  $s(t)$  is some function. Note the dependency on space and time via  $(t - x/c)$ , which denotes that it is a travelling wave. The sign in the argument is depending on which direction the wave travelling in.

Basic principles of the seismic method - TU Delft OCW

Seismic Impedance As any geophysical method, seismic reflection imaging is sensitive to only a specific physical property of the rock This property is called acoustic impedance:  $I = \rho v$  (density times acoustic velocity) Seismic reflection amplitude is proportional to the relative impedance contrast across a contact of two layers:  $\frac{I_2 - I_1}{I_2 + I_1}$

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This law is utilized in the seismic reflection method. It states that "the angle of incidence is equal to the angle of reflection". In case of  $I=0$ , the ratio of the reflected energy of P-wave,  $E_r$ , to the incident energy,  $E_i$ , is given by:  $\frac{E_r}{E_i} = 0$  The square root of the above relationship is called Reflection Coefficient,  $R$ .

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The general principle of seismic reflection is to send elastic waves (using an energy source such as dynamite explosion or Vibroseis) into the Earth, where each layer within the Earth reflects a portion of the wave's energy back and allows the rest to refract through. These reflected energy waves are recorded over a predetermined time period (called the record length) by receivers that detect the motion of the ground in which they are placed.

#### *Reflection seismology - Wikipedia*

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